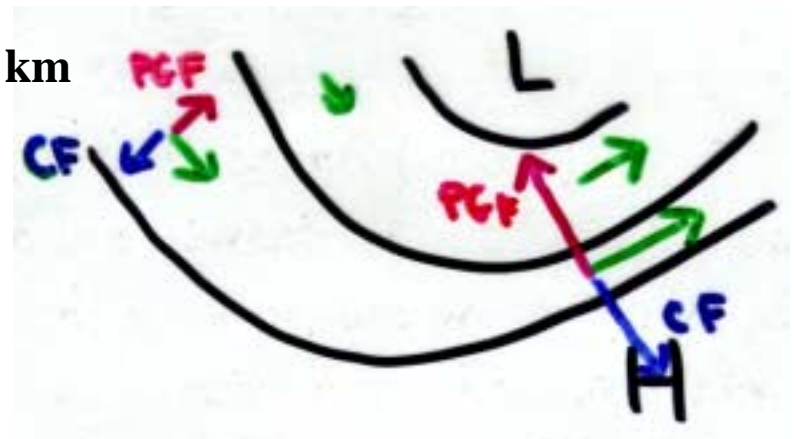


Lecture 12. Global Wind Systems

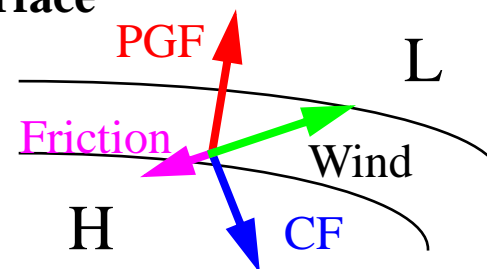
Review of Coriolis Force

- We see a Coriolis ‘force’ because wind is measured with respect to the rotating Earth -- an accelerating reference frame.
- Coriolis ‘force’ pushes moving objects to their right (in NH) or left (in SH).
- Above the lowest 1 km of the atmosphere, away from the Equator, the wind blows *along* isobars with high pressure on the right (NH), approximately balancing PGF and Coriolis forces.
- In the lowest 1 km or so, surface friction causes wind to spiral toward low pressure.

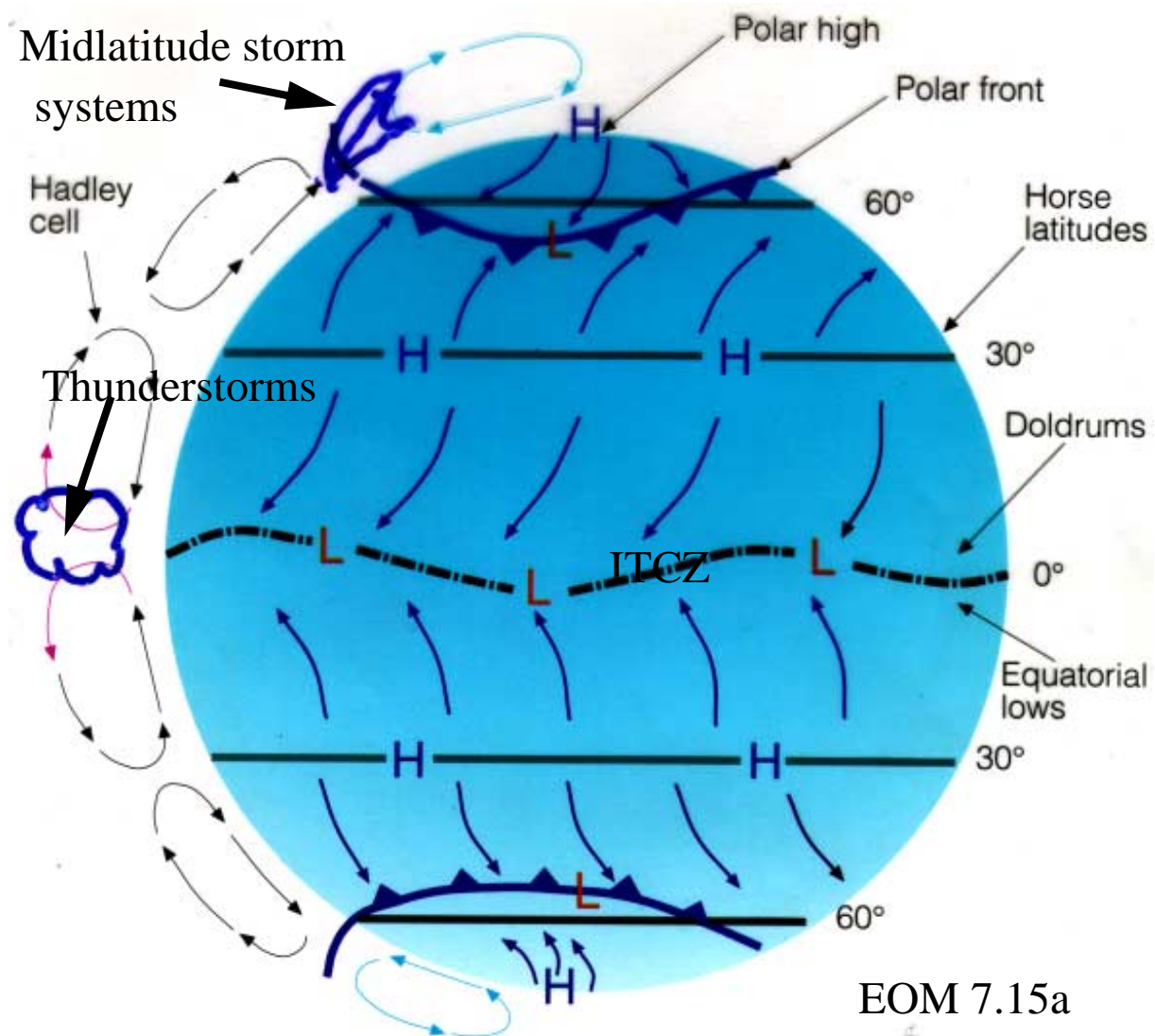
Above 1 km



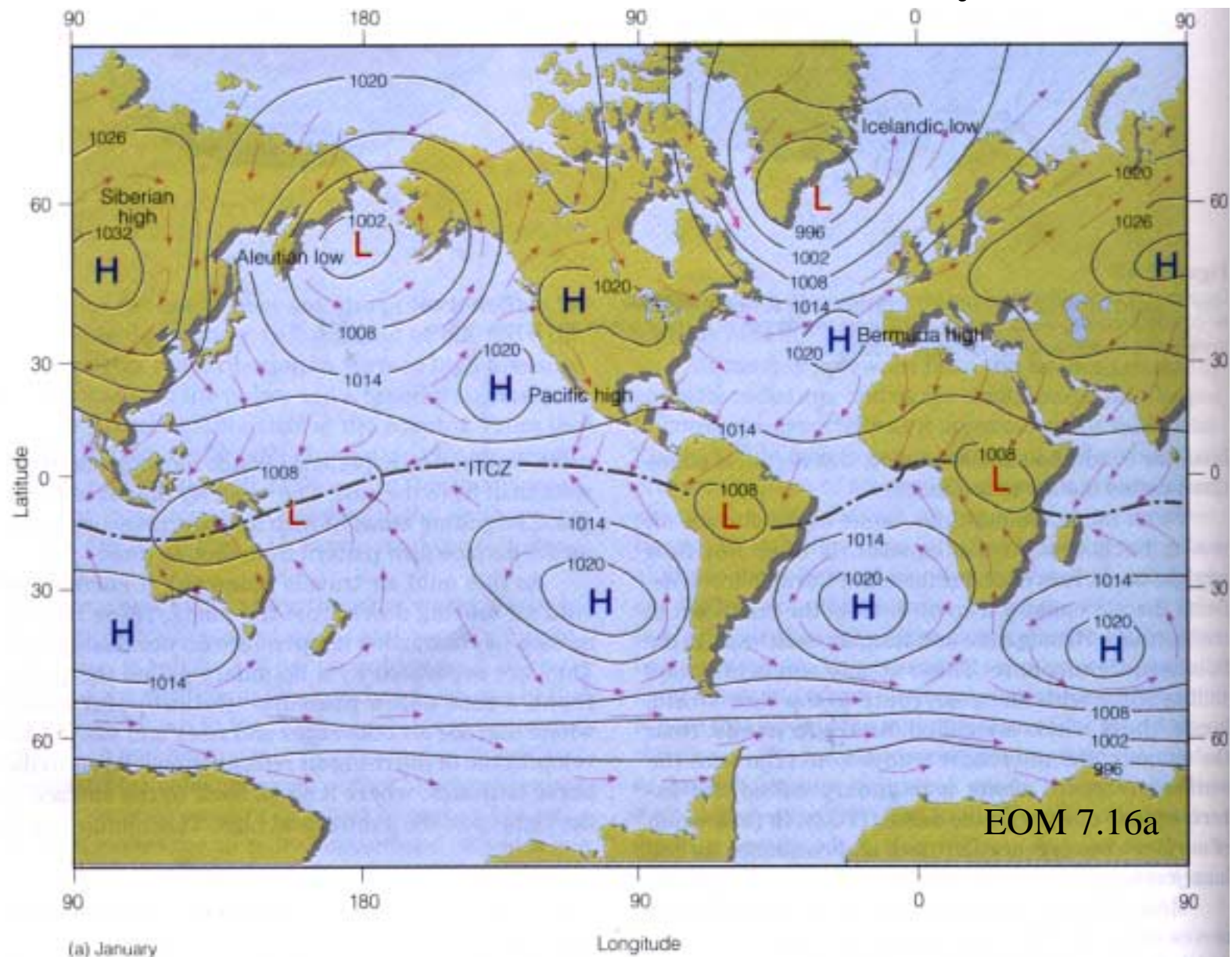
Near surface



Idealized General Circulation of Atmosphere



Mean Sea Level Pressure - January

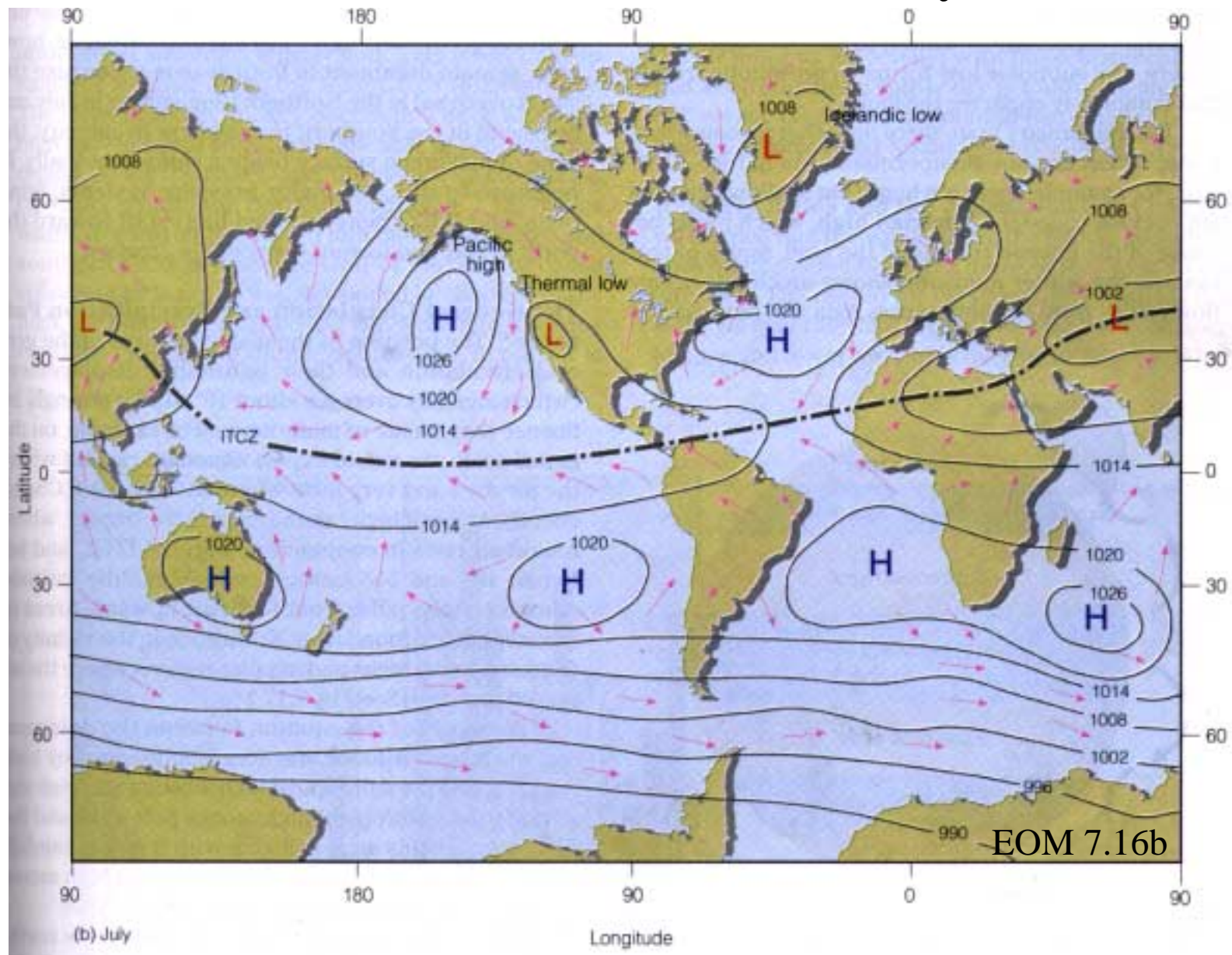


EOM 7.16a

Figure 7.16

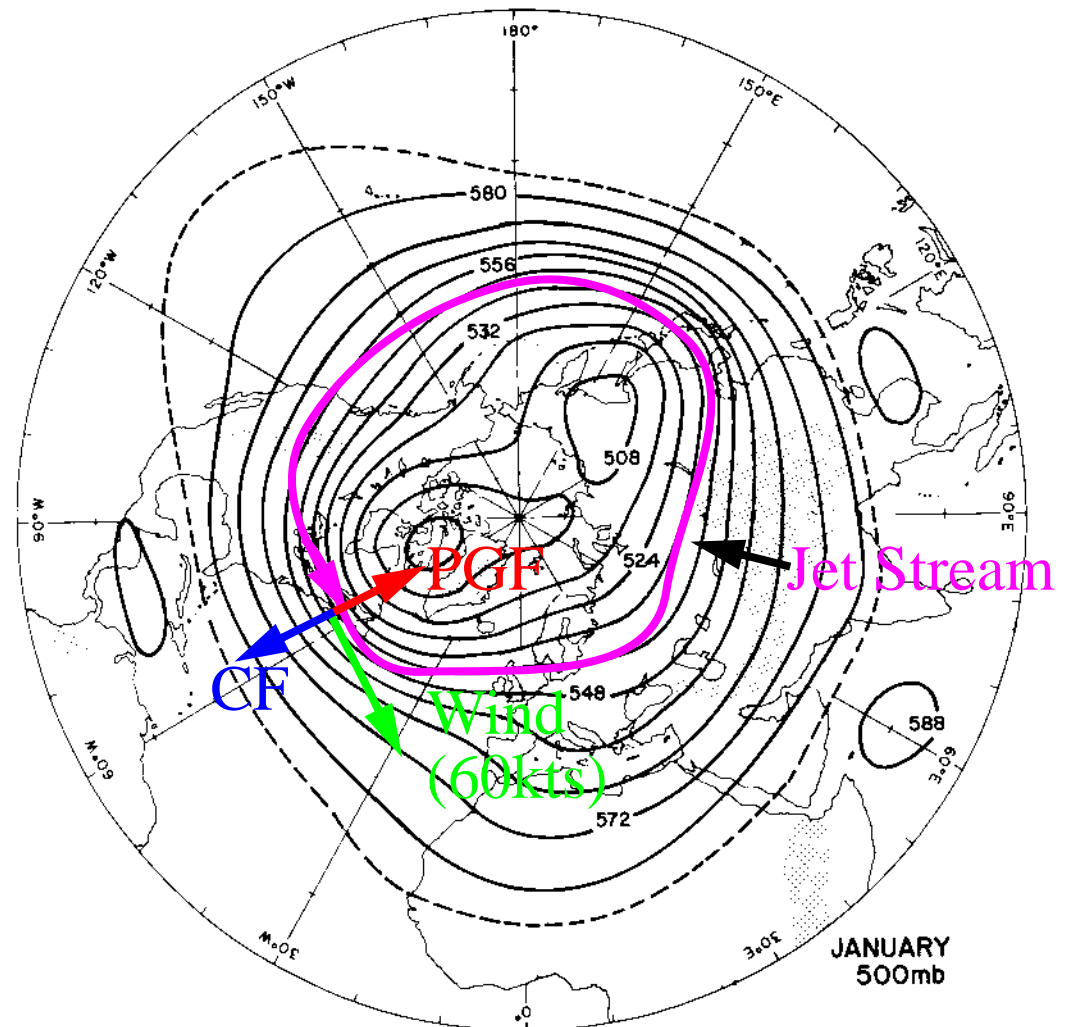
Average sea level pressure distribution and surface wind-flow patterns for January (a) and for July (b). The heavy dashed line represents the position of the ITCZ (Inter-Tropical Convergence Zone)

Mean Sea Level Pressure - July

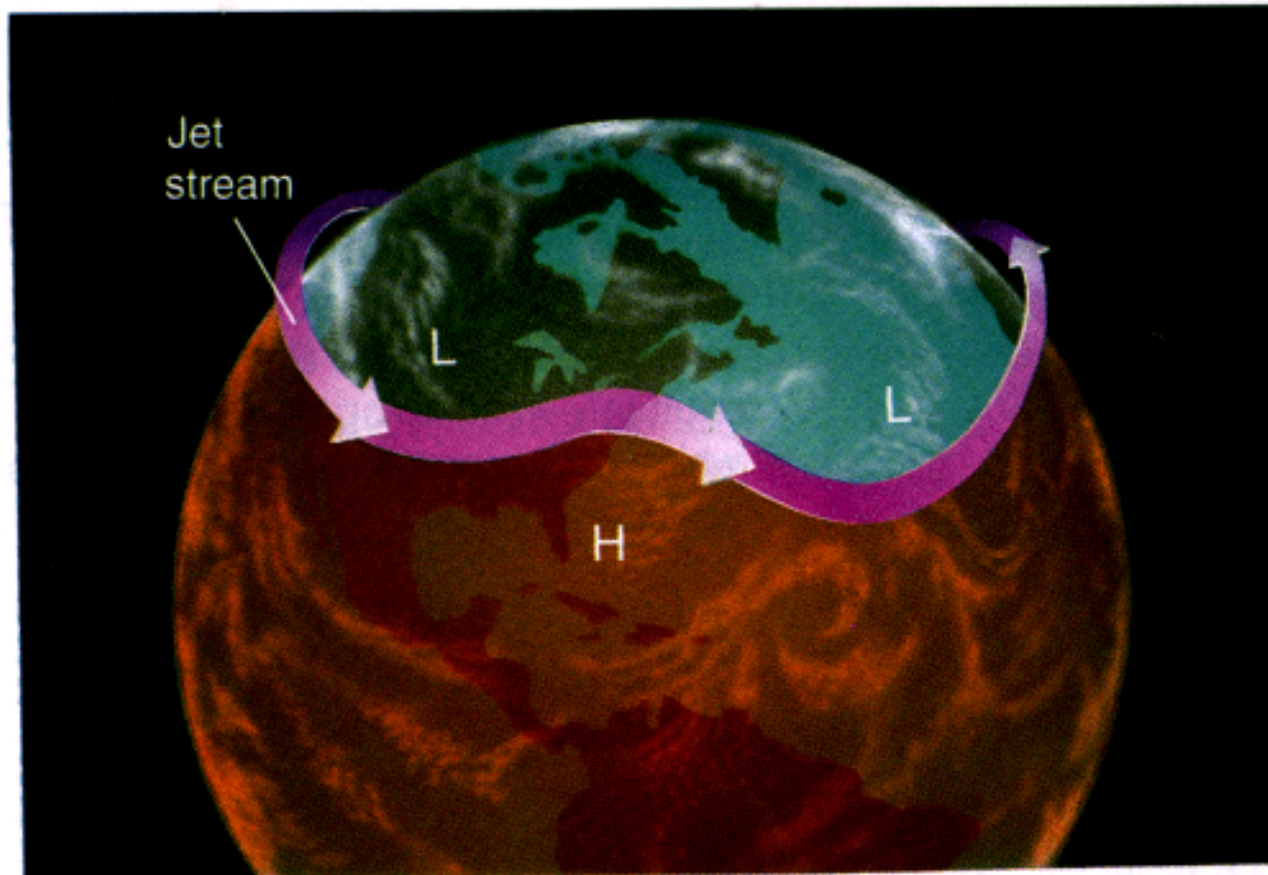


Height of 500 mb surface - January

- 1 km higher at equator than poles.
- Much larger pressure differences than at surface drive jet stream (mean winds to 60 kts).



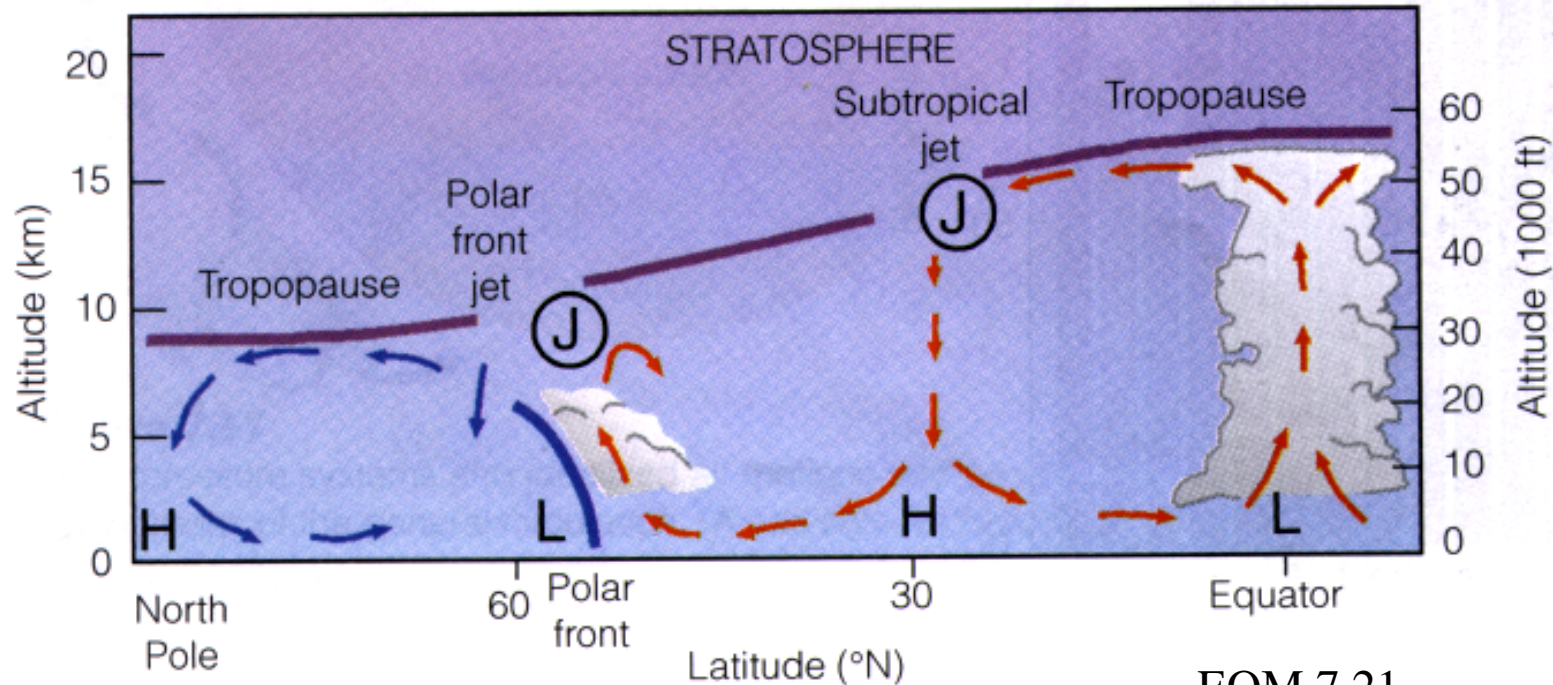
Jet Streams



EOM **Figure 7.20**

A jet stream is a swiftly flowing current of air that moves in a wavy west-to-east direction. It forms along a boundary where colder air lies to the north and warmer air to the south.

Typical height of jet streams and tropopause



EOM 7.21

- Note tropopause is higher near equator than near poles
- Jet stream winds are strongest near the tropopause
- Sometimes the jet stream separates into two wind maxima, a *polar* and a *subtropical* jet, but at most longitudes and times, a single jet is seen.

Formation of Jet Streams

- Temperature is colder (air denser) toward the poles in troposphere (below tropopause, typically about 300 mb).
 - Pressure drops faster with height in cold air
- ⇒ Hor. pressure gradient grows up to 300 mb
- ⇒ Jet stream maximum winds are at 300 mb.

