Time Differencing: Physically Insignificant Fast Waves

Dale Durran

Department of Atmospheric Sciences University of Washington

May 18, 2011



A question of efficiency

Suppose

- ▶ The vertical grid spacing is 300 m.
- ▶ The speed of sound is 300 m s $^{-1}$.

A question of efficiency

Suppose

- The vertical grid spacing is 300 m.
- ▶ The speed of sound is 300 m s $^{-1}$.
- ▶ A fully explicit time differencing approximation to the full compressible governing equations will require the time step to be no larger than roughly 1 second.

- Filtering them from the governing equations.
 - Boussinesq, Anelastic or Pseudo-incompressible equations
 - Numerical solution via the projection method

- Filtering them from the governing equations.
 - ▶ Boussinesq, Anelastic or Pseudo-incompressible equations
 - Numerical solution via the projection method
- Choosing a numerical method that treats the terms responsible for sound wave propagation efficiently.

- Filtering them from the governing equations.
 - ▶ Boussinesq, Anelastic or Pseudo-incompressible equations
 - Numerical solution via the projection method
- Choosing a numerical method that treats the terms responsible for sound wave propagation efficiently.
 - Semi-implicit methods

- Filtering them from the governing equations.
 - ▶ Boussinesq, Anelastic or Pseudo-incompressible equations
 - Numerical solution via the projection method
- Choosing a numerical method that treats the terms responsible for sound wave propagation efficiently.
 - Semi-implicit methods
 - Splitting methods

- Filtering them from the governing equations.
 - ▶ Boussinesq, Anelastic or Pseudo-incompressible equations
 - Numerical solution via the projection method
- Choosing a numerical method that treats the terms responsible for sound wave propagation efficiently.
 - Semi-implicit methods
 - Splitting methods
 - The approximation of (high frequency) sound waves is wildly inaccurate

Outline

The Projection Method Boussinesq context Solution procedure

The Semi-Implicit Method

Large time steps and poor accuracy The oscillation equation The shallow-water equations Stratified flow

Splitting Methods

Complete operator splitting Partial operator splitting

Summary



Boussinesq equations

$$\frac{\partial \mathbf{v}}{\partial t} + \nabla P = \mathbf{F} \equiv -\mathbf{v} \cdot \nabla \mathbf{v} + b\mathbf{k},$$
$$\frac{db}{dt} + N^2 w = 0,$$
$$\nabla \cdot \mathbf{v} = 0.$$

Boussinesq equations

$$\frac{\partial \mathbf{v}}{\partial t} + \nabla P = \mathbf{F} \equiv -\mathbf{v} \cdot \nabla \mathbf{v} + b\mathbf{k},$$
$$\frac{db}{dt} + N^2 w = 0,$$
$$\nabla \cdot \mathbf{v} = 0.$$

Here

- P is the Boussinesq pressure potential
- b is buoyancy
- ▶ N is the Brunt-Väisälä (buoyancy) frequency

$$\int_{t^n}^{t^{n+1}} \frac{\partial \mathbf{v}}{\partial t} dt = -\int_{t^n}^{t^{n+1}} \nabla P dt + \int_{t^n}^{t^{n+1}} \mathbf{F}(\mathbf{v}, b) dt$$

$$\int_{t^n}^{t^{n+1}} \frac{\partial \mathbf{v}}{\partial t} dt = -\int_{t^n}^{t^{n+1}} \nabla P dt + \int_{t^n}^{t^{n+1}} \mathbf{F}(\mathbf{v}, b) dt$$

Define \tilde{P}^{n+1} such that

$$\Delta t \, \nabla \tilde{P}^{n+1} = \int_{t^n}^{t^{n+1}} \nabla P \, dt;$$

$$\int_{t^n}^{t^{n+1}} \frac{\partial \mathbf{v}}{\partial t} dt = -\int_{t^n}^{t^{n+1}} \nabla P dt + \int_{t^n}^{t^{n+1}} \mathbf{F}(\mathbf{v}, b) dt$$

Define \tilde{P}^{n+1} such that

$$\Delta t \, \nabla \tilde{P}^{n+1} = \int_{t^n}^{t^{n+1}} \nabla P \, dt;$$

 $ilde{P}$ is not the pressure at any particular time.

$$\int_{t^n}^{t^{n+1}} \frac{\partial \mathbf{v}}{\partial t} dt = -\int_{t^n}^{t^{n+1}} \nabla P dt + \int_{t^n}^{t^{n+1}} \mathbf{F}(\mathbf{v}, b) dt$$

Define \tilde{P}^{n+1} such that

$$\Delta t \, \nabla \tilde{P}^{n+1} = \int_{t^n}^{t^{n+1}} \nabla P \, dt;$$

 \tilde{P} is not the pressure at any particular time.

$$\mathbf{v}^{n+1} - \mathbf{v}^n = -\Delta t \, \nabla \tilde{P}^{n+1} + \int_{t^n}^{t^{n+1}} \mathbf{F}(\mathbf{v}, b) \, dt.$$

Update the velocities in two steps

1) Compute $\tilde{\mathbf{v}}$, defined such that

$$\tilde{\mathbf{v}} = \mathbf{v}^n + \int_{t^n}^{t^{n+1}} \mathbf{F}(\mathbf{v}, b) dt.$$

Update the velocities in two steps

1) Compute $\tilde{\mathbf{v}}$, defined such that

$$\tilde{\mathbf{v}} = \mathbf{v}^n + \int_{t^n}^{t^{n+1}} \mathbf{F}(\mathbf{v}, b) dt.$$

Recall

$$\mathbf{v}^{n+1} - \mathbf{v}^n = -\Delta t \, \nabla \tilde{P}^{n+1} + \int_{t^n}^{t^{n+1}} \mathbf{F}(\mathbf{v}, b) \, dt.$$

Update the velocities in two steps

1) Compute $\tilde{\mathbf{v}}$, defined such that

$$\tilde{\mathbf{v}} = \mathbf{v}^n + \int_{t^n}^{t^{n+1}} \mathbf{F}(\mathbf{v}, b) dt.$$

Recall

$$\mathbf{v}^{n+1} - \mathbf{v}^n = -\Delta t \, \nabla \tilde{P}^{n+1} + \int_{t^n}^{t^{n+1}} \mathbf{F}(\mathbf{v}, b) \, dt.$$

2) After evaluating \tilde{P} , complete the update

$$\mathbf{v}^{n+1} = \tilde{\mathbf{v}} - \Delta t \, \nabla \tilde{P}^{n+1}.$$

Evaluating \tilde{P}^{\parallel}

From
$$\nabla \cdot \mathbf{v}^{n+1} = 0$$
 and

$$\mathbf{v}^{n+1} = \tilde{\mathbf{v}} - \Delta t \, \nabla \tilde{P}^{n+1},$$

Evaluating \tilde{P}

From $\nabla \cdot \mathbf{v}^{n+1} = 0$ and

$$\mathbf{v}^{n+1} = \tilde{\mathbf{v}} - \Delta t \, \nabla \tilde{P}^{n+1},$$

 \tilde{P}^{n+1} satisfies

$$\nabla^2 \tilde{P}^{n+1} = \frac{\nabla \cdot \tilde{\mathbf{v}}}{\Delta t}.$$

Evaluating \tilde{P}

From $\nabla \cdot \mathbf{v}^{n+1} = 0$ and

$$\mathbf{v}^{n+1} = \tilde{\mathbf{v}} - \Delta t \, \nabla \tilde{P}^{n+1},$$

 \tilde{P}^{n+1} satisfies

$$abla^2 \tilde{P}^{n+1} = rac{
abla \cdot \tilde{\mathbf{v}}}{\Delta t}.$$

Specification of boundary conditions for this Poisson equation can be an nontrivial detail.

Interpretation as a projection

The operation

$$\mathbf{v}^{n+1} = \tilde{\mathbf{v}} - \Delta t \, \nabla \tilde{P}^{n+1},$$

projects the partially updated velocities $\tilde{\mathbf{v}}$ onto the subspace of non-divergent velocity fields.

Interpretation as a projection

The operation

$$\mathbf{v}^{n+1} = \tilde{\mathbf{v}} - \Delta t \, \nabla \tilde{P}^{n+1},$$

projects the partially updated velocities $\tilde{\mathbf{v}}$ onto the subspace of non-divergent velocity fields.

This projection has no influence on the time truncation error.

Interpretation as a projection

The operation

$$\mathbf{v}^{n+1} = \tilde{\mathbf{v}} - \Delta t \, \nabla \tilde{P}^{n+1},$$

projects the partially updated velocities $\tilde{\mathbf{v}}$ onto the subspace of non-divergent velocity fields.

This projection has no influence on the time truncation error. Similar methods can also be used to integrate related sound-proof systems

- ► The anelastic equations
- ► The pseudo-incompressible equations

Large time steps and poor accuracy The oscillation equation The shallow-water equations Stratified flow

Outline

The Projection Method

Solution procedure

The Semi-Implicit Method

Large time steps and poor accuracy
The oscillation equation
The shallow-water equations
Stratified flow

Splitting Methods

Complete operator splitting Partial operator splitting

Summary



Semi-discrete approximations to the advection equation

Suppose the time derivative in

$$\frac{\partial \psi}{\partial t} + c \frac{\partial \psi}{\partial x} = 0$$

is approximated by leapfrog time differencing as

$$\frac{\phi^{n+1} - \phi^{n-1}}{2\Delta t} + c \left(\frac{\partial \phi}{\partial x}\right)^n = 0.$$

Leapfrog phase speed

Wave solutions of the form

$$\phi^n(x) = e^{i(kx - \omega n \Delta t)}$$

have phase speed

$$c_{\mathrm{lf}} = \frac{\omega}{k} = \frac{\mathrm{arcsin}(ck\Delta t)}{k\Delta t}.$$

Leapfrog phase speed

Wave solutions of the form

$$\phi^n(x) = e^{i(kx - \omega n \Delta t)}$$

have phase speed

$$c_{
m lf} = rac{\omega}{k} = rac{{
m arcsin}(ck\Delta t)}{k\Delta t}.$$

Stability requires

- $\blacktriangleright \omega$ real
- $ightharpoonup |ck\Delta t| < 1$

Trapezoidal phase speed

Using trapezoidal time differencing (over $2\Delta t$)

$$\frac{\phi^{n+1} - \phi^{n-1}}{2\Delta t} + \frac{c}{2} \left[\left(\frac{\partial \phi}{\partial x} \right)^{n+1} + \left(\frac{\partial \phi}{\partial x} \right)^{n-1} \right] = 0$$

Trapezoidal phase speed

Using trapezoidal time differencing (over $2\Delta t$)

$$\frac{\phi^{n+1} - \phi^{n-1}}{2\Delta t} + \frac{c}{2} \left[\left(\frac{\partial \phi}{\partial x} \right)^{n+1} + \left(\frac{\partial \phi}{\partial x} \right)^{n-1} \right] = 0$$

Solutions have phase speed

$$c_{
m t} = rac{\omega}{k} = rac{{
m arctan}(ck\Delta t)}{k\Delta t}.$$

Trapezoidal phase speed

Using trapezoidal time differencing (over $2\Delta t$)

$$\frac{\phi^{n+1} - \phi^{n-1}}{2\Delta t} + \frac{c}{2} \left[\left(\frac{\partial \phi}{\partial x} \right)^{n+1} + \left(\frac{\partial \phi}{\partial x} \right)^{n-1} \right] = 0$$

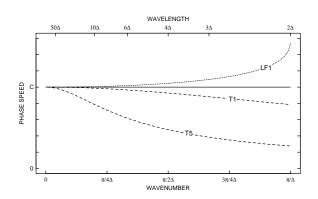
Solutions have phase speed

$$c_{
m t} = rac{\omega}{k} = rac{{
m arctan}(ck\Delta t)}{k\Delta t}.$$

- Unconditionally stable
- ▶ Phase speed reduced relative to that of the leapfrog solution by the factor $\cos(\omega \Delta t)$



Phase speed as a function of wavelength



Phase speed of leapfrog (dotted) and $2\Delta t$ -trapezoidal (dashed) approximations to the advection equation when

 $c\Delta t/\Delta x=1/\pi$ (LF1 and T1), and for the trapezoidal solution when $c\Delta t/\Delta x=5/\pi$ (T5).



A prototype ODE

$$\frac{d\psi}{dt} + i\omega_H \psi + i\omega_L \psi = 0$$

Suppose the high frequency forcing exceeds that of the low frequency forcing $|\omega_L|<|\omega_H|$.

A prototype ODE

$$\frac{d\psi}{dt} + i\omega_H \psi + i\omega_L \psi = 0$$

Suppose the high frequency forcing exceeds that of the low frequency forcing $|\omega_L| < |\omega_H|$.

Approximate as

$$\frac{\phi^{n+1} - \phi^{n-1}}{2\Delta t} + i\omega_H \left(\frac{\phi^{n+1} + \phi^{n-1}}{2}\right) + i\omega_L \phi^n = 0$$

Stability condition

Solutions of the form $\exp(-i\omega n\Delta t)$ exist provided

$$\sin(\omega \Delta t) = \omega_H \Delta t \cos(\omega \Delta t) + \omega_L \Delta t. \tag{1}$$

Stability condition

Solutions of the form $\exp(-i\omega n\Delta t)$ exist provided

$$\sin(\omega \Delta t) = \omega_H \Delta t \cos(\omega \Delta t) + \omega_L \Delta t. \tag{1}$$

Stable when

$$(\omega_L \Delta t)^2 \leq 1 + (\omega_H \Delta t)^2.$$

▶ Always satisfied when $|\omega_L| < |\omega_H|$.

Stability condition

Solutions of the form $\exp(-i\omega n\Delta t)$ exist provided

$$\sin(\omega \Delta t) = \omega_H \Delta t \cos(\omega \Delta t) + \omega_L \Delta t. \tag{1}$$

Stable when

$$(\omega_L \Delta t)^2 \leq 1 + (\omega_H \Delta t)^2$$
.

- ▶ Always satisfied when $|\omega_L| < |\omega_H|$.
- ▶ Always satisfied when $|\omega_L \Delta t| < 1$.
 - Same criterion is obtained if $\omega_H = 0$
 - ▶ The terms approximated implicitly have no impact on stability



Linearized one-dimensional shallow-water system

$$\begin{split} &\frac{\partial u}{\partial t} + U \frac{\partial u}{\partial x} + g \frac{\partial \eta}{\partial x} = 0, \\ &\frac{\partial \eta}{\partial t} + U \frac{\partial \eta}{\partial x} + H \frac{\partial u}{\partial x} = 0, \end{split}$$

- \triangleright U and u(x,t) are the mean and perturbation fluid velocity
- ▶ H and $\eta(x,t)$ are the mean and perturbation fluid depth

Numerical operator notation

Finite difference over interval $n\Delta t$

$$\delta_{nt}f(t) = \frac{f(t + n\Delta t/2) - f(t - n\Delta t/2)}{n\Delta t}$$

Numerical operator notation

Finite difference over interval $n\Delta t$

$$\delta_{nt}f(t) = \frac{f(t + n\Delta t/2) - f(t - n\Delta t/2)}{n\Delta t}$$

Averaging operator over interval $n\Delta t$

$$\langle f(t) \rangle^{nt} = \frac{f(t + n\Delta t/2) + f(t - n\Delta t/2)}{2}$$

Semi-discrete approximation

Semi-implicit approximation

$$\delta_{2t}u^n + U\frac{du^n}{dx} + g\left\langle\frac{d\eta^n}{dx}\right\rangle^{2t} = 0,$$

$$\delta_{2t}\eta^n + U \frac{d\eta^n}{dx} + H \left\langle \frac{du^n}{dx} \right\rangle^{2t} = 0.$$

Semi-discrete approximation

Semi-implicit approximation

$$\delta_{2t}u^n + U\frac{du^n}{dx} + g\left\langle\frac{d\eta^n}{dx}\right\rangle^{2t} = 0,$$

$$\delta_{2t}\eta^n + U \frac{d\eta^n}{dx} + H \left\langle \frac{du^n}{dx} \right\rangle^{2t} = 0.$$

Advection is leapfrog.

Semi-discrete approximation

Semi-implicit approximation

$$\delta_{2t}u^n + U\frac{du^n}{dx} + g\left\langle\frac{d\eta^n}{dx}\right\rangle^{2t} = 0,$$

$$\delta_{2t}\eta^n + U\frac{d\eta^n}{dx} + H\left\langle \frac{du^n}{dx} \right\rangle^{2t} = 0.$$

- Advection is leapfrog.
- ▶ Pressure gradient and divergence is trapezoidal over $2\Delta t$.

Semi-discrete dispersion relation

$$\sin \omega \Delta t = Uk\Delta t \pm \sqrt{gH}k\Delta t \cos \omega \Delta t$$

Semi-discrete dispersion relation

$$\sin \omega \Delta t = Uk\Delta t \pm \sqrt{gH}k\Delta t \cos \omega \Delta t$$

▶ Same form as (1) for the oscillation equation

Semi-discrete dispersion relation

$$\sin \omega \Delta t = Uk\Delta t \pm \sqrt{gH}k\Delta t \cos \omega \Delta t$$

- ▶ Same form as (1) for the oscillation equation
- ▶ Always stable if flow is subcritical $(|U| < \sqrt{gH})$

Nonlinear one-dimensional shallow-water system

$$\frac{\partial u}{\partial t} + (U + u)\frac{\partial u}{\partial x} + g\frac{\partial \eta}{\partial x} = 0,$$

$$\frac{\partial \eta}{\partial t} + (U + u)\frac{\partial \eta}{\partial x} + (H + \eta)\frac{\partial u}{\partial x} = 0.$$

Nonlinear one-dimensional shallow-water system

$$\begin{split} \frac{\partial u}{\partial t} + (U + u) \frac{\partial u}{\partial x} + g \frac{\partial \eta}{\partial x} &= 0, \\ \frac{\partial \eta}{\partial t} + (U + u) \frac{\partial \eta}{\partial x} + (H + \eta) \frac{\partial u}{\partial x} &= 0. \end{split}$$

Semi-discrete approximation is

$$\delta_{2t}u^n + (U + u^n)\frac{du^n}{dx} + g\left\langle\frac{d\eta^n}{dx}\right\rangle^{2t} = 0,$$

$$\delta_{2t}\eta^n + (U+u^n)\frac{d\eta^n}{dx} + H\left\langle\frac{du^n}{dx}\right\rangle^{2t} + \eta^n\frac{du^n}{dx} = 0.$$

Nonlinear one-dimensional shallow-water system

$$\begin{split} \frac{\partial u}{\partial t} + (U + u) \frac{\partial u}{\partial x} + g \frac{\partial \eta}{\partial x} &= 0, \\ \frac{\partial \eta}{\partial t} + (U + u) \frac{\partial \eta}{\partial x} + (H + \eta) \frac{\partial u}{\partial x} &= 0. \end{split}$$

Semi-discrete approximation is

$$\delta_{2t}u^n + (U + u^n)\frac{du^n}{dx} + g\left\langle\frac{d\eta^n}{dx}\right\rangle^{2t} = 0,$$

$$\delta_{2t}\eta^n + (U + u^n)\frac{d\eta^n}{dx} + H\left\langle\frac{du^n}{dx}\right\rangle^{2t} + \eta^n\frac{du^n}{dx} = 0.$$

Nonlinear part of the divergence is leapfrog.



Large time steps and poor accuracy The oscillation equation The shallow-water equations Stratified flow

Influence on stability

How does explicit differencing of a portion of the divergence influence stability?

Large time steps and poor accuracy The oscillation equation The shallow-water equations Stratified flow

Influence on stability

How does explicit differencing of a portion of the divergence influence stability?

Linearize about a basic state at rest with non-zero perturbation displacement $\overline{\eta}$.

Influence on stability

How does explicit differencing of a portion of the divergence influence stability?

- Linearize about a basic state at rest with non-zero perturbation displacement $\overline{\eta}$.
- ▶ Stability requires $|\overline{\eta}| < H$, which seems easy to satisfy.

Influence on stability

How does explicit differencing of a portion of the divergence influence stability?

- Linearize about a basic state at rest with non-zero perturbation displacement $\overline{\eta}$.
- ▶ Stability requires $|\overline{\eta}| < H$, which seems easy to satisfy.
- ▶ Local phase speed $\sqrt{g(H+\overline{\eta})}$ cannot exceed the reference phase speed \sqrt{gH} by more than $\sqrt{2}$.

Hydrostatic internal gravity waves propagate at the same speed as a shallow-water gravity wave of the appropriate equivalent depth.

Hydrostatic internal gravity waves propagate at the same speed as a shallow-water gravity wave of the appropriate equivalent depth.

The horizontal phase speed of hydrostatic internal gravity waves is proportional to the Brunt-Väisälä frequency.

▶ Compare \sqrt{gH} and N/ℓ .

Hydrostatic internal gravity waves propagate at the same speed as a shallow-water gravity wave of the appropriate equivalent depth.

The horizontal phase speed of hydrostatic internal gravity waves is proportional to the Brunt-Väisälä frequency.

► Compare \sqrt{gH} and N/ℓ .

The ratio of the local mean N in the polar regions of the earth's atmosphere to that in mid-latitudes can easily exceed $\sqrt{2}$.

Hydrostatic internal gravity waves propagate at the same speed as a shallow-water gravity wave of the appropriate equivalent depth.

The horizontal phase speed of hydrostatic internal gravity waves is proportional to the Brunt-Väisälä frequency.

▶ Compare \sqrt{gH} and N/ℓ .

The ratio of the local mean N in the polar regions of the earth's atmosphere to that in mid-latitudes can easily exceed $\sqrt{2}$.

Choose the reference-state stratification as isothermal.



Compressible Boussinesq system

Boussinesq system with prognostic equation for pressure retained.

$$\begin{split} \frac{d\mathbf{v}}{dt} + \nabla P &= b\mathbf{k}, \\ \frac{db}{dt} + N^2 w &= 0, \\ \frac{dP}{dt} + c_s^2 \nabla \cdot \mathbf{v} &= 0. \end{split}$$

(Simple system supporting both sound and gravity waves.)

Semi-implicit formulation

$$\mathbf{v}^{n+1} + \Delta t \nabla P^{n+1} = \hat{\mathbf{v}} \equiv \mathbf{v}^{n-1} - \Delta t \left[\nabla P^{n-1} - 2b^n \mathbf{k} + 2\mathbf{v}^n \cdot \nabla \mathbf{v}^n \right]$$
$$P^{n+1} + c_s^2 \Delta t \nabla \cdot \mathbf{v}^{n+1} = h \equiv P^{n-1} - c_s^2 \Delta t \left[\nabla \cdot \mathbf{v}^{n-1} + 2\mathbf{v}^n \cdot \nabla P^n \right].$$

Semi-implicit formulation

$$\mathbf{v}^{n+1} + \Delta t \nabla P^{n+1} = \hat{\mathbf{v}} \equiv \mathbf{v}^{n-1} - \Delta t \left[\nabla P^{n-1} - 2b^n \mathbf{k} + 2\mathbf{v}^n \cdot \nabla \mathbf{v}^n \right]$$
$$P^{n+1} + c_s^2 \Delta t \nabla \cdot \mathbf{v}^{n+1} = h \equiv P^{n-1} - c_s^2 \Delta t \left[\nabla \cdot \mathbf{v}^{n-1} + 2\mathbf{v}^n \cdot \nabla P^n \right].$$

Substituting the divergence of the first into the second yields

$$\nabla^2 P^{n+1} - \frac{P^{n+1}}{(c_s \Delta t)^2} = \frac{\nabla \cdot \hat{\mathbf{v}}}{\Delta t} - \frac{h}{(c_s \Delta t)^2}$$

Semi-implicit formulation

$$\mathbf{v}^{n+1} + \Delta t \nabla P^{n+1} = \hat{\mathbf{v}} \equiv \mathbf{v}^{n-1} - \Delta t \left[\nabla P^{n-1} - 2b^n \mathbf{k} + 2\mathbf{v}^n \cdot \nabla \mathbf{v}^n \right]$$
$$P^{n+1} + c_s^2 \Delta t \nabla \cdot \mathbf{v}^{n+1} = h \equiv P^{n-1} - c_s^2 \Delta t \left[\nabla \cdot \mathbf{v}^{n-1} + 2\mathbf{v}^n \cdot \nabla P^n \right].$$

Substituting the divergence of the first into the second yields

$$\nabla^2 P^{n+1} - \frac{P^{n+1}}{(c_s \Delta t)^2} = \frac{\nabla \cdot \hat{\mathbf{v}}}{\Delta t} - \frac{h}{(c_s \Delta t)^2}$$

Compare with corresponding result from the projection method

$$\nabla^2 \tilde{p}^{n+1} = \frac{\nabla \cdot \tilde{\mathbf{v}}}{\Delta t}.$$



Outline

The Projection Method

Boussinesq context Solution procedure

The Semi-Implicit Method

Large time steps and poor accuracy
The oscillation equation
The shallow-water equations
Stratified flow

Splitting Methods

Complete operator splitting Partial operator splitting

Summary



Two fractional steps

A partial differential equation of the general form

$$\frac{\partial \psi}{\partial t} + \mathcal{L}_1(\psi) + \mathcal{L}_2(\psi) = 0,$$

may be approximated using two fractional steps as

$$\phi^s = \mathcal{F}_1(\Delta t)\phi^n,$$

$$\phi^{n+1} = \mathcal{F}_2(\Delta t)\phi^s$$
.

Two fractional steps

A partial differential equation of the general form

$$\frac{\partial \psi}{\partial t} + \mathcal{L}_1(\psi) + \mathcal{L}_2(\psi) = 0,$$

may be approximated using two fractional steps as

$$\phi^{s} = \mathcal{F}_{1}(\Delta t)\phi^{n},$$

$$\phi^{n+1} = \mathcal{F}_{2}(\Delta t)\phi^{s}.$$

▶ $\mathcal{F}_i(\Delta t)\phi^n$ approximates the action of the exact operator \mathcal{L}_i mapping $\psi(t_n)$ to $\psi(t_{n+1})$.

Two fractional steps

A partial differential equation of the general form

$$\frac{\partial \psi}{\partial t} + \mathcal{L}_1(\psi) + \mathcal{L}_2(\psi) = 0,$$

may be approximated using two fractional steps as

$$\phi^{s} = \mathcal{F}_{1}(\Delta t)\phi^{n},$$

$$\phi^{n+1} = \mathcal{F}_{2}(\Delta t)\phi^{s}.$$

- ▶ $\mathcal{F}_i(\Delta t)\phi^n$ approximates the action of the exact operator \mathcal{L}_i mapping $\psi(t_n)$ to $\psi(t_{n+1})$.
- ▶ If \mathcal{L}_i is a time-independent linear operator

$$\psi(t+\Delta t) = \exp(\Delta t \mathcal{L}_i)\psi(t) = \left(I + \Delta t \mathcal{L}_i + \frac{(\Delta t)^2}{2}\mathcal{L}_i^2 + \ldots\right)\psi(t).$$

Unequal time steps

Suppose \mathcal{L}_2 contains those terms responsible for the propagation of waves moving a factor of M faster than those propagated by \mathcal{L}_1 .

Unequal time steps

Suppose \mathcal{L}_2 contains those terms responsible for the propagation of waves moving a factor of M faster than those propagated by \mathcal{L}_1 .

A first-order splitting (unless the operators commute) is

$$\phi^{n+1} = \left[\mathcal{F}_2(\Delta t/M)\right]^M \mathcal{F}_1(\Delta t)\phi^n.$$

Unequal time steps

Suppose \mathcal{L}_2 contains those terms responsible for the propagation of waves moving a factor of M faster than those propagated by \mathcal{L}_1 .

A first-order splitting (unless the operators commute) is

$$\phi^{n+1} = \left[\mathcal{F}_2(\Delta t/M) \right]^M \mathcal{F}_1(\Delta t) \phi^n.$$

A second-order (Strang) splitting is

$$\phi^{n+1} = \left[\mathcal{F}_2(2\Delta t/M)\right]^{(M/2)} \mathcal{F}_1(\Delta t) \left[\mathcal{F}_2(2\Delta t/M)\right]^{(M/2)} \phi^n.$$

A test case

Compressible 2D Boussinesq system

$$\frac{\partial \mathbf{r}}{\partial t} + \mathcal{L}_1(\mathbf{r}) + \mathcal{L}_2(\mathbf{r}) = \mathbf{F}(x, z, t),$$

where $\mathbf{r} = \begin{pmatrix} u & w & b & P \end{pmatrix}^T$ and a nondivergent forcing is applied to the momentum equations

$$\mathbf{F} = \left(\begin{array}{ccc} -\frac{\partial \Psi}{\partial z} & \frac{\partial \Psi}{\partial x} & \mathbf{0} & \mathbf{0} \end{array} \right)^T.$$

Here Ψ is a streamfunction with compact support.

Advection step

Use linearly 3rd-order Runge-Kutta scheme and a large time step to advance the advection operator

$$\mathcal{L}_1 = \left(egin{array}{cccc} \mathbf{v} \cdot
abla & 0 & 0 & 0 & 0 \ 0 & \mathbf{v} \cdot
abla & 0 & 0 & 0 \ 0 & 0 & 0 & \mathbf{v} \cdot
abla \end{array}
ight)$$

Pressure gradient, divergence and buoyancy

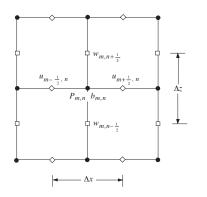
Use a small time step and forward-backward differencing, except for terms related to vertical propagation of sound waves, which are trapezoidal to improve stability.

Pressure gradient, divergence and buoyancy

Use a small time step and forward-backward differencing, except for terms related to vertical propagation of sound waves, which are trapezoidal to improve stability.

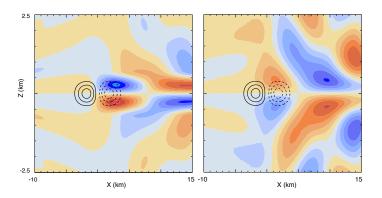
$$\begin{split} \frac{u^{m+1}-u^m}{\Delta\tau} + \frac{\partial P^m}{\partial x} &= -\frac{\partial \Psi^{m+\frac{1}{2}}}{\partial z}, \\ \frac{w^{m+1}-w^m}{\Delta\tau} + \frac{\partial}{\partial z} \left(\frac{P^{m+1}+P^m}{2}\right) - b^m &= \frac{\partial \Psi^{m+\frac{1}{2}}}{\partial x}, \\ \frac{b^{m+1}-b^m}{\Delta\tau} + N^2 w^{m+1} &= 0, \\ \frac{P^{m+1}-P^m}{\Delta\tau} + c_s^2 \frac{\partial u^{m+1}}{\partial x} + c_s^2 \frac{\partial}{\partial z} \left(\frac{w^{m+1}+w^m}{2}\right) &= 0, \end{split}$$

More details

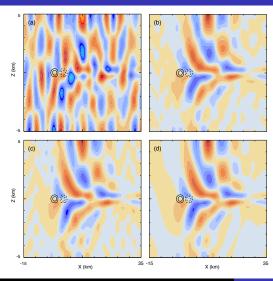


- Centered 2nd-order spatial differences on staggered mesh
- Approximation to $\exp(\Delta t \mathcal{L}_1)$ is stable for $|U|\Delta t/\Delta x < 1.73$.
- Approximation to $\exp(\Delta \tau \mathcal{L}_2)$ is stable and non-damping for $\max(c_s/\Delta x, N)\Delta \tau < 1$.

Spatially uniform U

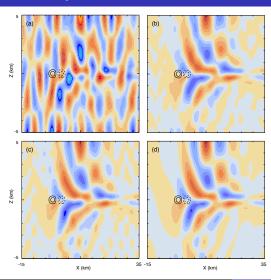


Contours of (a) U + u and Ψ ; (b) P and Ψ . No zero contour is drawn. Minor tick marks indicate the location of the P points on the numerical grid. Only the central portion of the domain is shown.



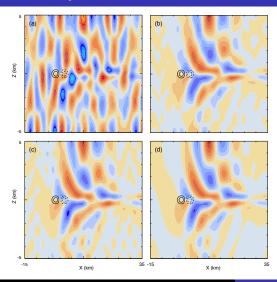
Contours of P and Ψ . Tick marks every 20 grid intervals

(a) same Δt and $\Delta \tau$



Contours of P and Ψ . Tick marks every 20 grid intervals

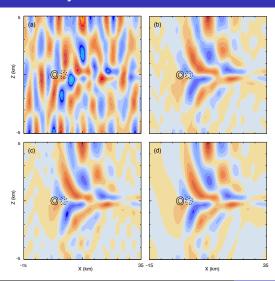
- (a) same Δt and $\Delta \tau$
- (b) Both Δt and $\Delta \tau$ halved



Contours of P and Ψ . Tick marks every 20 grid

intervals

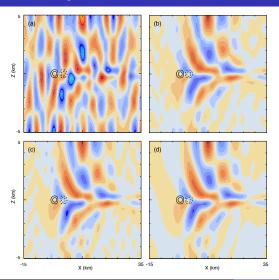
- (a) same Δt and $\Delta \tau$
- (b) Both Δt and $\Delta \tau$ halved
- (c) Only Δt halved



Contours of P and Ψ . Tick marks every 20 grid

intervals

- (a) same Δt and $\Delta \tau$
- (b) Both Δt and $\Delta \tau$ halved
- (c) Only Δt halved
- (d) Partially split: same Δt and $\Delta \tau$



Contours of P and Ψ . Tick marks every 20 grid intervals

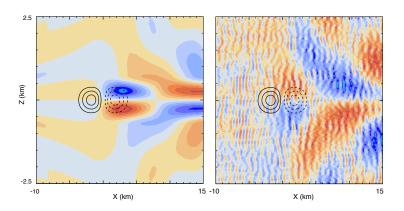
- (a) same Δt and $\Delta \tau$
- (b) Both Δt and Δau halved
- (c) Only Δt halved
- (d) Partially split: same Δt and $\Delta \tau$

In all cases U + u remains correct at this time.

Include a piece of the advective forcing on every small time step.

$$\begin{split} \frac{u^{m+1}-u^m}{\Delta\tau} + \frac{\partial P^m}{\partial x} &= -U\frac{\partial u^n}{\partial x} - w^n\frac{\partial U}{\partial z},\\ \frac{w^{m+1}-w^m}{\Delta\tau} + \frac{\partial}{\partial z}\left(\frac{P^{m+1}+P^m}{2}\right) - b^m &= -U\frac{\partial w^n}{\partial x},\\ \frac{b^{m+1}-b^m}{\Delta\tau} + N^2w^{m+1} &= -U\frac{\partial b^n}{\partial x},\\ \frac{P^{m+1}-P^m}{\Delta\tau} + c_s^2\frac{\partial u^{m+1}}{\partial x} + c_s^2\frac{\partial}{\partial z}\left(\frac{w^{m+1}+w^m}{2}\right) &= -U\frac{\partial P^n}{\partial x}. \end{split}$$

Spatially uniform U - revisited



Contours of (a) U + u and Ψ ; (b) P (contoured using twice the interval as in the completely split case) and Ψ .



Fixing the partially split method

The partially split method is unstable.

Fixing the partially split method

The partially split method is unstable.

Yet it is widely used in many mesoscale models.

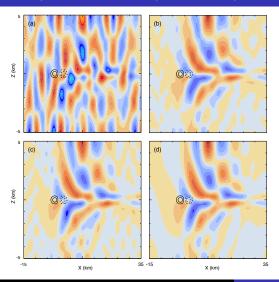
Fixing the partially split method

The partially split method is unstable.

Yet it is widely used in many mesoscale models.

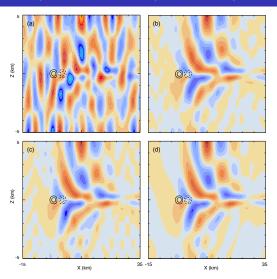
It can be stabilized using

- ► The Asselin time filter (used to prevent time-splitting in leapfrog integrations) (Tatsumi, 1983)
- ▶ Damping the velocity divergence every small time step (Skamarock and Klemp, 1992)



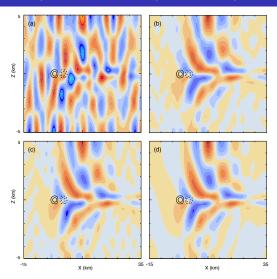
Contours of P and Ψ . Tick marks every 20 grid intervals. $\frac{\partial U}{\partial z} \neq 0$.

(a) Complete splitting, pressure noisy.



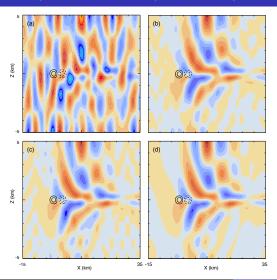
Contours of P and Ψ . Tick marks every 20 grid intervals. $\frac{\partial U}{\partial z} \neq 0$.

- (a) Complete splitting, pressure noisy.
- (b) Both Δt and Δau halved



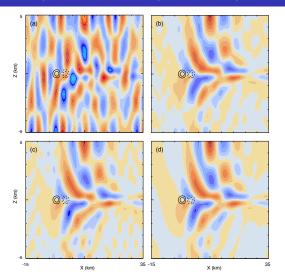
Contours of P and Ψ . Tick marks every 20 grid intervals. $\frac{\partial U}{\partial z} \neq 0$.

- (a) Complete splitting, pressure noisy.
- (b) Both Δt and Δau halved
- (c) Only Δt halved



Contours of P and Ψ . Tick marks every 20 grid intervals. $\frac{\partial U}{\partial z} \neq 0$.

- (a) Complete splitting, pressure noisy.
- (b) Both Δt and $\Delta \tau$ halved
- (c) Only Δt halved
- (d) Partially splitting: same Δt and $\Delta \tau$ as in (a)



Contours of P and Ψ . Tick marks every 20 grid intervals. $\frac{\partial U}{\partial z} \neq 0$.

- (a) Complete splitting, pressure noisy.
- (b) Both Δt and Δau halved
- (c) Only Δt halved
- (d) Partially splitting: same Δt and $\Delta \tau$ as in (a)

Divergence damping does not significantly improve the completely split solutions.

Outline

The Projection Method

Boussinesq context Solution procedure

The Semi-Implicit Method

Large time steps and poor accuracy
The oscillation equation
The shallow-water equations
Stratified flow

Splitting Methods

Complete operator splitting Partial operator splitting

Summary



The pressure P and divergence δ in the compressible Boussinesq system satisfy

$$\begin{split} \frac{\partial P}{\partial t} + c_s^2 \delta &= -\mathbf{v} \cdot \nabla P, \\ \frac{\partial \delta}{\partial t} + \nabla^2 P &= -\nabla \cdot (\mathbf{v} \cdot \nabla \mathbf{v}) + \frac{\partial b}{\partial z}. \end{split}$$

The pressure P and divergence δ in the compressible Boussinesq system satisfy

$$\begin{split} \frac{\partial P}{\partial t} + c_s^2 \delta &= -\mathbf{v} \cdot \nabla P, \\ \frac{\partial \delta}{\partial t} + \nabla^2 P &= -\nabla \cdot (\mathbf{v} \cdot \nabla \mathbf{v}) + \frac{\partial b}{\partial z}. \end{split}$$

Discard prognostic equation for P.

The pressure P and divergence δ in the compressible Boussinesq system satisfy

$$\frac{\partial P}{\partial t} + c_s^2 \delta = -\mathbf{v} \cdot \nabla P,$$

$$\frac{\partial \delta}{\partial t} + \nabla^2 P = -\nabla \cdot (\mathbf{v} \cdot \nabla \mathbf{v}) + \frac{\partial b}{\partial z}.$$

- Discard prognostic equation for P.
- ▶ Close 2nd equation by settting $\delta = 0$, and solve Poisson equation for P.

The pressure P and divergence δ in the compressible Boussinesq system satisfy

$$\frac{\partial P}{\partial t} + c_s^2 \delta = -\mathbf{v} \cdot \nabla P,$$

$$\frac{\partial \delta}{\partial t} + \nabla^2 P = -\nabla \cdot (\mathbf{v} \cdot \nabla \mathbf{v}) + \frac{\partial b}{\partial z}.$$

- Discard prognostic equation for P.
- ▶ Close 2nd equation by settting $\delta = 0$, and solve Poisson equation for P.
- Wide range of choices for integrating remaining terms



$$\begin{split} \frac{\partial P}{\partial t} + c_s^2 \delta &= - \mathbf{v} \cdot \nabla P, \\ \frac{\partial \delta}{\partial t} + \nabla^2 P &= - \nabla \cdot \left(\mathbf{v} \cdot \nabla \mathbf{v} \right) + \frac{\partial b}{\partial z}. \end{split}$$

$$\begin{split} \frac{\partial P}{\partial t} + c_s^2 \delta &= -\mathbf{v} \cdot \nabla P, \\ \frac{\partial \delta}{\partial t} + \nabla^2 P &= -\nabla \cdot \left(\mathbf{v} \cdot \nabla \mathbf{v}\right) + \frac{\partial b}{\partial z}. \end{split}$$

Uses trapezoidal differencing for the left sides

$$\begin{split} \frac{\partial P}{\partial t} + c_s^2 \delta &= -\mathbf{v} \cdot \nabla P, \\ \frac{\partial \delta}{\partial t} + \nabla^2 P &= -\nabla \cdot \left(\mathbf{v} \cdot \nabla \mathbf{v}\right) + \frac{\partial b}{\partial z}. \end{split}$$

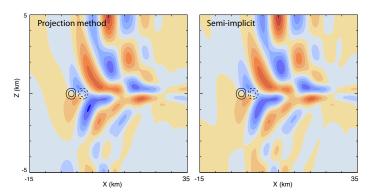
- Uses trapezoidal differencing for the left sides
- Solve Helmholtz equation for P.

$$\begin{split} \frac{\partial P}{\partial t} + c_s^2 \delta &= -\mathbf{v} \cdot \nabla P, \\ \frac{\partial \delta}{\partial t} + \nabla^2 P &= -\nabla \cdot \left(\mathbf{v} \cdot \nabla \mathbf{v}\right) + \frac{\partial b}{\partial z}. \end{split}$$

- Uses trapezoidal differencing for the left sides
- Solve Helmholtz equation for P.
- Leapfrog differencing for integrating remaining terms

Test case with shear

Contours of P and Ψ as in previous.



Both the Projection and Semi-Implicit methods yield results similar to that obtained with partial splitting and divergence damping.

$$\frac{\partial^2 \delta}{\partial t^2} - c_s^2 \nabla^2 \delta = \frac{\partial^2 b}{\partial t \partial z}.$$

During the small-time-step cycle, the divergence satisfies

$$\frac{\partial^2 \delta}{\partial t^2} - c_s^2 \nabla^2 \delta = \frac{\partial^2 b}{\partial t \partial z}.$$

► The divergence propagates as an undamped wave, subject to some modification by buoyancy forces.

$$\frac{\partial^2 \delta}{\partial t^2} - c_s^2 \nabla^2 \delta = \frac{\partial^2 b}{\partial t \partial z}.$$

- ► The divergence propagates as an undamped wave, subject to some modification by buoyancy forces.
- ▶ Divergence generated on the large time step is added to the initial condition for δ at the start of each cycle, and builds throughout the integration.

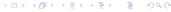
$$\frac{\partial^2 \delta}{\partial t^2} - c_s^2 \nabla^2 \delta = \frac{\partial^2 b}{\partial t \partial z}.$$

- ► The divergence propagates as an undamped wave, subject to some modification by buoyancy forces.
- ▶ Divergence generated on the large time step is added to the initial condition for δ at the start of each cycle, and builds throughout the integration.
- Advection by constant U is a Galilean translation that does not generate divergence.



$$\frac{\partial^2 \delta}{\partial t^2} - c_s^2 \nabla^2 \delta = \frac{\partial^2 b}{\partial t \partial z}.$$

- ► The divergence propagates as an undamped wave, subject to some modification by buoyancy forces.
- ▶ Divergence generated on the large time step is added to the initial condition for δ at the start of each cycle, and builds throughout the integration.
- Advection by constant U is a Galilean translation that does not generate divergence.
- Divergence is generated when there is vertical shear.



► The divergence is almost zero at the start of the first small step.

- ► The divergence is almost zero at the start of the first small step.
- lacktriangle Only small changes in δ are forced during a single large time step.

- ► The divergence is almost zero at the start of the first small step.
- lacktriangle Only small changes in δ are forced during a single large time step.
- ▶ The divergence forced over each small step is similar to that which would appear in an integration of the full compressible system provided the amplitudes of the sound waves are small compared to those of the slower modes.

- ► The divergence is almost zero at the start of the first small step.
- \blacktriangleright Only small changes in δ are forced during a single large time step.
- ▶ The divergence forced over each small step is similar to that which would appear in an integration of the full compressible system provided the amplitudes of the sound waves are small compared to those of the slower modes.
- Divergence damping provides a feedback keeping sound-wave amplitudes small.



Iterative solvers versus small steps

▶ The Poisson equations arising in the Boussinesq system could be solved by iteration.

Iterative solvers versus small steps

- ► The Poisson equations arising in the Boussinesq system could be solved by iteration.
- ▶ The inaccuracy of the completely split method suggests the small step cycle does not constitute an iteration process capable of arriving at the pressure that correctly projects the evolving velocity field onto the nondivergent (or anelastic equivalent) subspace.

Iterative solvers versus small steps

- ► The Poisson equations arising in the Boussinesq system could be solved by iteration.
- ➤ The inaccuracy of the completely split method suggests the small step cycle does not constitute an iteration process capable of arriving at the pressure that correctly projects the evolving velocity field onto the nondivergent (or anelastic equivalent) subspace.
- ▶ The partial splitting method succeeds because the forcing for the divergence stays close to that actually associated with the propagation of the slow modes.

References

Durran, D.R., 1999: Numerical Methods for Wave Equations in Geophysical Fluid Dynamics, Springer. (Sections 7.1—7.4)

Skamarock, W.C. and J.B. Klemp, 1992: The stability of time-split numerical methods for the hydrostatic and nonhydrostatic elastic equations. *Mon. Wea. Rev.*, **120**, 2109-2171.