## **Homework Set 2 (Monin-Obukhov theory)**

1. (Adapted from Arya, p. 180) The following measurements of mean wind and potential temperature were taken around noon during the 1968 Kansas field program:

z (m)	2	4	8	16	32	
<i>u</i> (m s <sup>-1</sup> )	5.81	6.70	7.49	8.14	8.66	
$\theta(K)$	307.20	306.65	306.28	305.88	305.62	

We wish to calculate the surface fluxes of heat and moisture using the Monin-Obuhkov relations  $\phi_h = \phi_m^2 = \{1 - 16z/L\}^{-1/2}$  for an unstable surface layer. To proceed:

(a) Calculate the gradients of u and  $\theta$  by differencing between successive heights. In the surface layer, the profiles tend to vary roughly logarithmically with height, so it is better to difference using  $\ln(z)$  as the height coordinate. Note that  $du/dz = z^{-1} \frac{du}{d}(\ln z)$  and similarly for  $\theta$ . Numbering the levels 1 (2 m) to 5 (32 m), the estimated gradient of u between levels 1 and 2 would be

$$\left(\frac{\Delta u}{\Delta z}\right)_{21} = \frac{1}{z_m} \left(\frac{\Delta u}{\Delta \ln z}\right)_{21} = \frac{1}{z_m} \left(\frac{u_2 - u_1}{\ln z_2 - \ln z_1}\right)$$

at a height  $z_m$  such that  $\ln z_m = (\ln z_1 + \ln z_2)/2 = (\ln 2 + \ln 4)/2$ , i. e. at  $z_m = 2.82$  m. This works out to a gradient  $(\Delta u/\Delta z)_{12} = 0.45$  s<sup>-1</sup>. Neglecting virtual effects on buoyancy, use your calculated gradients to find Ri vs. z.

- (b) From this data, estimate the Obukhov length L (note that Ri = z/L in an unstable surface layer).
- (c) Using the data from the lowest two heights and your L from (b), calculate the friction velocity, the sensible heat flux, and the surface roughness length  $z_0$ . Does the implied roughness length seem appropriate for a field of wheat stubble? Take the air density =  $1.2 \text{ kg m}^{-3}$  and  $C_p = 10^3 \text{ J kg}^{-1} \text{ K}^{-1}$ . Can we also estimate the thermal roughness length  $z_T$  from the given data?
- 2. An oceanographic research ship is stationed 100 km off the California coast. It is measuring a wind speed of 10 m s<sup>-1</sup> and an air temperature of 286.9 K at a height of 10 m above sea level (i.e. 287 K if adiabatically displaced to the sea surface). The ocean surface temperature is also 287 K. Neglect virtual effects, so we have a neutral surface layer.
  - (a) Using the bulk aerodynamic approach, with  $C_{DN} = (0.75 + 0.067u_{10}) \times 10^{-3}$ , find the surface stress and the friction velocity (use the same reference air density as before)?
  - (b) What is the roughness length  $z_0$  (use Charnock's formula)?
  - (c) If you based  $C_{DN}$  on this  $z_0$ , rather than the bulk formula of (a), how different would the result be (this checks the consistency of the two approaches).
  - (d) The saturation mixing ratio for salt water at the sea-surface is 9.7 g kg<sup>-1</sup>, while the mixing ratio at 10 m elevation is 7.4 g kg<sup>-1</sup>. Using a bulk formula with  $C_{qN} = 1.3 \times 10^{-3}$ , calculate the surface latent heat flux.
  - (e) An airplane flies overhead at 30 m elevation. Assuming this is still within the surface layer, what mean wind speed, temperature and mixing ratio would the aircraft measure?
- 3. Now the ship moves over a coastal upwelling zone, where the 10 m wind and air temperature remain as above but the SST is cooled 3 K to 284 K.

- (a) Starting with the neutral bulk formula of (a) and  $C_{HN} = 1.3 \times 10^{-3}$ , make a first guess at the sensible heat flux and Obukhov length L. This won't be exact, but will be good enough for our purposes.
- (b) Using this to estimate z/L at the 10 m measurement height, use Garratt Fig. 3.7 (shown in the Lecture 5 notes; use the  $z/z_0 = 10^5$  curve) to estimate by what percentage  $C_D$  is decreased from  $C_{DN}$ ? Based on this estimate, by what percentage is the surface wind stress changed by moving over the colder SST?