

Flow in a channel; Kelvin Waves (CR6.2, G11 10.4)

In a channel, $v=0$ at $y=0, W$. Governing LSW: $y=0 \text{ --- } y=W$
 with $v = \text{Re } \tilde{v}(y) e^{i(kx - \omega t)}$:

$$\frac{\partial}{\partial t} \left\{ -\frac{\partial^2}{\partial t^2} - f^2 + gH_0 \left(\frac{\partial^2}{\partial x^2} + \frac{\partial^2}{\partial y^2} \right) \right\} v = 0 \quad y=0 \text{ --- } y=W \quad x \rightarrow$$

$$\Rightarrow -i\omega \left\{ gH_0 \frac{d^2 \tilde{v}}{dy^2} + (\omega^2 - f^2 - gH_0 k^2) \tilde{v} \right\} = 0 \quad \tilde{v}(0) = \tilde{v}(W) = 0.$$

Possibilities: (I) $\omega=0$ (leads to geostrophic balanced mode with $\tilde{v} = \frac{\partial \tilde{h}}{\partial y} = 0$ at $y=0, W$)

$$(II) \quad gH_0 \frac{d^2 \tilde{v}}{dy^2} + (\omega^2 - f^2 - gH_0 k^2) \tilde{v} = 0$$

$$\tilde{v} = \sin l_n y, \quad \omega^2 = f^2 + gH_0 (k^2 + l_n^2), \quad l_n = \frac{n\pi}{W} \quad (\text{Standing Poincaré wave})$$

(III) $\tilde{v} = 0$ everywhere!

$$\Rightarrow \frac{\partial v}{\partial t} = -g \frac{\partial h}{\partial x}$$

$$\frac{\partial v}{\partial t} + f v = -g \frac{\partial h}{\partial y}$$

Geostrophic balance in y

$$\frac{\partial h}{\partial t} + H \left(\frac{\partial v}{\partial x} + \frac{\partial v}{\partial y} \right) = 0$$

$$\begin{bmatrix} \eta \\ u \end{bmatrix} = \text{Re } e^{i(kx - \omega t)} \begin{bmatrix} \tilde{\eta} \\ \tilde{u} \end{bmatrix}(y)$$

$$\Rightarrow -i\omega \tilde{u} = -ikg \tilde{\eta} \quad \text{or } \tilde{u} = \frac{kg}{\omega} \tilde{\eta}$$

$$-i\omega \tilde{\eta} + ikH \tilde{u} = 0 \quad \Rightarrow -i\omega + ikH \frac{kg}{\omega} = 0 \Rightarrow \omega^2 = k^2 c_0^2$$

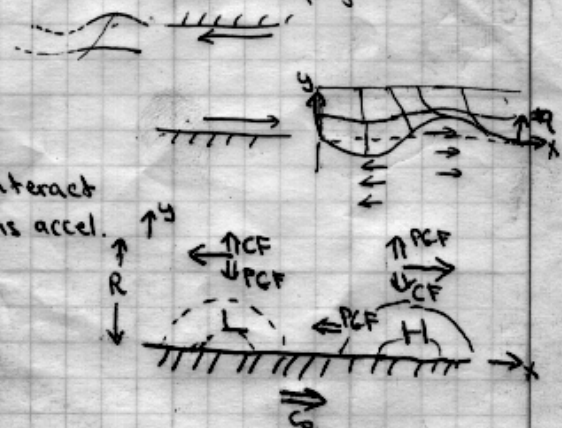
Two roots: $\omega = kc_0$ (R-travelling Kelvin wave), $\frac{d\tilde{\eta}}{dy} = -\frac{f}{c_0} \tilde{\eta} \Rightarrow \tilde{\eta} = \tilde{\eta} e^{-y/R}$
 $\omega = -kc_0$ (L-travelling Kelvin wave), $\frac{d\tilde{\eta}}{dy} = \frac{c_0}{f} \tilde{\eta} \Rightarrow \tilde{\eta} = \tilde{\eta} e^{y/R}$

Kelvin waves are non-dispersive + like gravity waves in propagation direction, supported by geostrophic balance in normal direction. Propagate with wall on their R. (if $f > 0$)

For a R-moving Kelvin wave, $\tilde{u} = \frac{g}{c_0} \tilde{\eta}$

is in phase with $\tilde{\eta}$.

Fluid "leans" on the wall to counteract Coriolis accel.



Equatorial Kelvin wave (Bill 11.5)

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A Kelvin wave in which the fluid in each hemisphere "leans" on the other hemisphere; equator takes place of wall

LSWE on β -plane

$$\frac{\partial u}{\partial t} + \beta y v = -g \frac{\partial \eta}{\partial x}$$

$$\frac{\partial v}{\partial t} + \beta y u = -g \frac{\partial \eta}{\partial y}$$

$$\frac{\partial \eta}{\partial t} + H \left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} \right) = 0$$

Look for mode with $v=0$, $\begin{pmatrix} u \\ \eta \end{pmatrix} = \begin{pmatrix} \hat{u}(y) \\ \hat{\eta}(y) \end{pmatrix} e^{i(kx - \omega t)}$

$$-i\omega \hat{u} = -ikg \hat{\eta} \Rightarrow \hat{u} = kg \hat{\eta} / \omega \text{ as before} \quad (1)$$

$$\beta y \hat{u} = -g \frac{\partial \hat{\eta}}{\partial y} \quad (\text{geostrophic balance in } y) \quad (2)$$

$$-i\omega \hat{\eta} + ikH \hat{u} = 0 \Rightarrow -i\omega + ikH \left\{ \frac{kg}{\omega} \right\} = 0 \text{ as before}$$

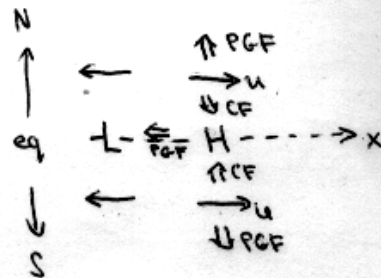
so $\omega = \pm c_0 k$, $\hat{u} = \pm g \hat{\eta} / c_0$

Plug into (2):

$$\pm \frac{\beta y \hat{\eta}}{c_0} = \pm g \frac{\partial \hat{\eta}}{\partial y}$$

$$\Rightarrow \frac{\partial \hat{\eta}}{\hat{\eta}} = \mp \frac{\beta y}{c_0} dy$$

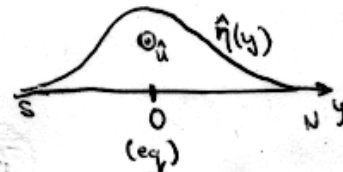
$$\hat{\eta} = \eta_0 e^{\mp \frac{\beta y^2}{2c_0}}$$



Only the top soln, $\frac{\omega}{k} = +c_0$, which is E-ward propagating is bold (and $\rightarrow 0$) as $y \rightarrow \pm \infty$, and is the equatorial Kelvin wave

Very similar to wall-trapped K-wave, but meridional scale is $R_{eq} = \left(\frac{c_0}{\beta} \right)^{\frac{1}{2}}$.

Oceanic equatorial Kelvin waves play a key role in the dynamics of El Niño



Atmospheric Kelvin waves are often coupled to enhanced tropical thunderstorm activity and rainfall, especially in the central Pacific ITCZ during boreal summer. The Madden-Julian Oscillation, though not really a Kelvin wave, also has related features (equatorial trapping, E-ward propagation) despite a slower propagation speed of $5-10 \frac{m}{s}$ in the Indian/West Pac ocean. Equatorial Kelvin waves propagating up into stratosphere help excite Quasi-Biennial Osc (QBO).