#### What is GFD?

ATM S 509/OCEAN 512
Geophysical Fluid Dynamics I
1-4-17

# Key aspects of geophysical flows

- Rotation
- Stratification

#### Also:

- Topography/bathymetry/ inhomogeneous surface
- Latent heating/other diabatic terms

- Development of equations for the atmosphere and ocean
  - Fluid properties of air and water
  - Ways to talk about stratification
  - Mathematical formulation of rotating reference frame

- Conservation equations
  - Energy
  - Vorticity
  - Potential vorticity

- Scaling arguments
  - To rule out particular terms and find dominant balances
  - Deriving simplified equations
  - Sometimes based on order of magnitudes
  - Sometimes formal asymptotic expansions in small nondimensional parameters

- Examination of transient and steady response to forcing
  - Waves
    - Linear theory, Fourier analysis

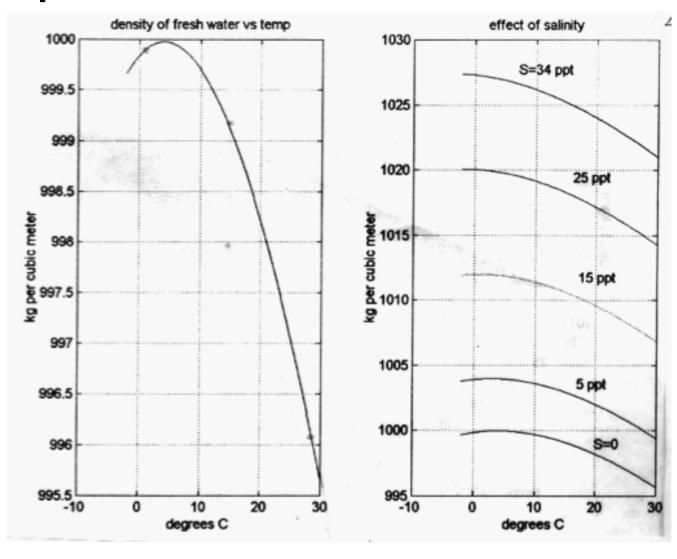
- Instabilities
  - Convective
  - Barotropic, baroclinic, inertial

- Idealized physical systems
  - To illustrate phenomena most succinctly and elegantly
- Lab experiments/numerical experiments

## **Equations of Motion**

- Conserved tracer
- Momentum equation
- Mass equation
- Equation of state

## Equation of state for seawater



## Equation of state for seawater

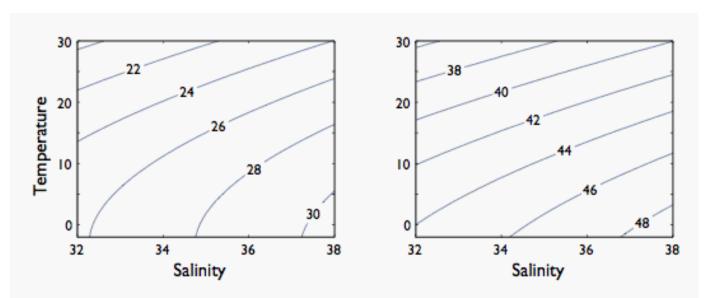


Fig. 1.3 A temperature-salinity diagram for seawater, calculated using an accurate empirical equation of state. Contours are (density -1000) kg m<sup>-3</sup>, and the temperature is potential temperature, which in the deep ocean may be less than *in situ* temperature by a degree or so (see Fig. 1.4). Left panel: at sea-level ( $p = 10^5$  Pa = 1000 mb). Right panel: at  $p = 4 \times 10^7$  Pa, a depth of about 4 km. Note that in both cases the contours are slightly convex.

## Equation of state for seawater

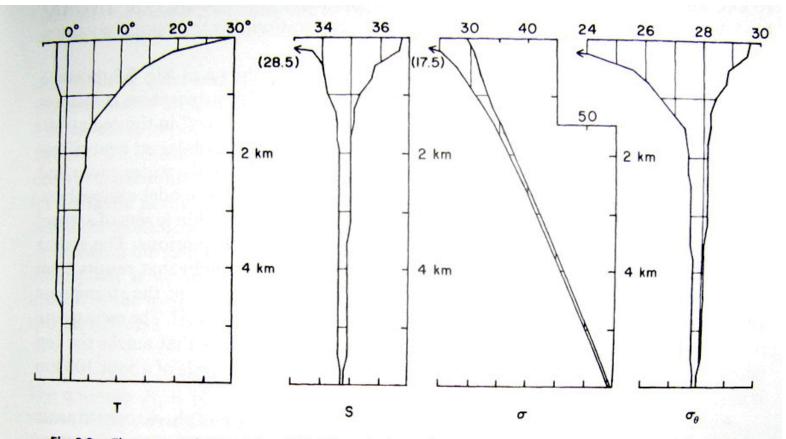
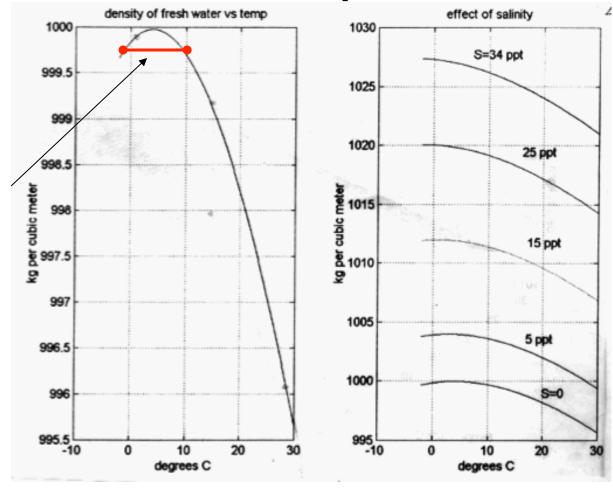


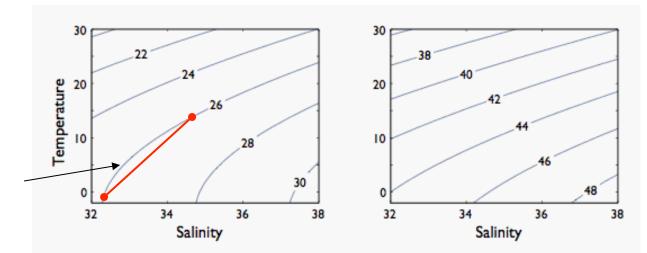
Fig. 3.2. The ranges of temperature T (in °C) and salinity S for 98% of the ocean as a function of depth [From Bryan and Cox (1972)], and the corresponding ranges of density  $\sigma$  and potential density  $\sigma_{\theta}$  (see Appendix 3).

## More on nonlinear eqn of state

Mix these two parcels



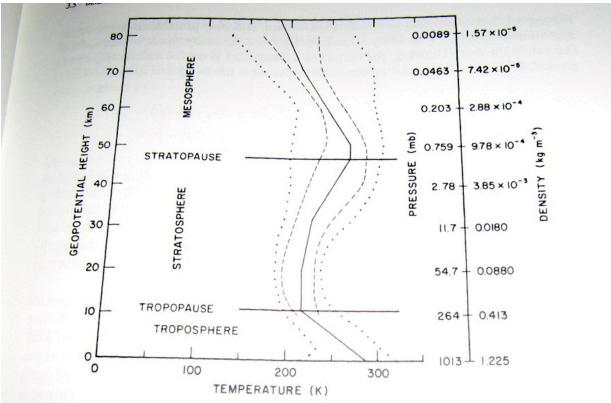
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Fig. 1.3 A temperature-salinity diagram for seawater, calculated using an accurate empirical equation of state. Contours are (density -1000) kg m $^{-3}$ , and the temperature is potential temperature, which in the deep ocean may be less than *in situ* temperature by a degree or so (see Fig. 1.4). Left panel: at sea-level ( $p=10^5\,\mathrm{Pa}=1000\,\mathrm{mb}$ ). Right panel: at  $p=4\times10^7\,\mathrm{Pa}$ , a depth of about  $4\,\mathrm{km}$ . Note that in both cases the contours are slightly convex.

## Density of Air vs Ocean



**Fig. 3.3.** Temperature variation with geopotential height for the U.S. Standard Atmosphere (solid line). This consists of straight-line segments with breaks at 11, 20, 32, 47, 51, and 71 km. The surface temperature is 15°C and the gradients, starting from the surface, are -6.5, 0. 1.0, 2.8, 0, -2.8, and -2.0 K km<sup>-1</sup>. The dashed line shows the lowest and highest monthly mean temperatures obtained for any location between equator and pole, whereas the dotted line shows estimates of the 1% maximum and minimum temperatures that occur during the warmest and coldest months, respectively, in the most extreme locations. The scale at the right gives pressures and densities at 10-km intervals for the standard profile. [From NOAA/NASA/USAF, 1976.]

Air is much less dense than water (3 orders of mag or more)

Also air density changes much more than sea water (orders of mag vs 2%)